

## **I. Abstract**

We explore the use of satellite observations to constrain the global NOx emissions. The GEOS-Chem chemistry and transport model, driven by assimilated NASA/GMAO GEOS-3 meteorological data (1979-2003), and its adjoint are applied for global inversions of SCIAMACHY tropospheric NO<sub>2</sub> columns. The targeted state vector contributing to tropospheric NO<sub>2</sub> column includes NOx emissions from fossil fuel, biomass burning, biofuel burning, soil, lightning, and NH<sub>3</sub> oxidation. We conducted a "pseudo inversion" of SCIAMACHY NO<sub>2</sub> data from November 2005 using GEOS-Chem results for November 2001, for which the GEOS-Chem adjoint is available (for GEOS-3). Our focus with this pseudo inversion is on demonstrating the utilities of the adjoint approach with satellite observations. Preliminary results from this inversion are presented here. The model overestimation of NOx over the Eastern US and Central Europe are improved by reducing significantly *a posteriori* fossil fuel emissions. The *a posteriori* biomass burning NOx emissions are higher over Siberia and Australia, while biofuel burning NOx contributions are lower over Eastern Europe and Asia. However, these changes may be largely attributed to temporal mismatches between model year (2001) vs. observations year (2005). The implementation of an adjoint for GEOS-Chem driven by GEOS-4 reanalysis (2003 and beyond) is ongoing.

## II. Introduction

Atmospheric nitrogen oxide radicals (NOx = NO + NO<sub>2</sub>) have profound influences on both tropospheric ozone, a key factor in air quality and climate change, and the hydroxyl radical, the primary oxidizing agent and scavenger of many species. Traditional bottom-up estimates of NOx emissions are highly uncertain due to extrapolation of high spatiotemporal variability of emission fluxes. Additional top-down constraints based on satellite observations are necessary to accurately represent the spatial and temporal variations in emissions. Here we derive top-down NOx emissions with SCIAMACHY tropospheric NO<sub>2</sub> column data using a variational approach with GEOS-Chem adjoint [Henze et al., 2007]. This method allows the use of finer spatial-temporal distribution of emissions and can account for nonlinear chemistry. It relies on the exact and efficient evaluation of the gradient of a cost function with respect to the control variables. The results from the adjoint inversion will be compared with those from a mass-balance method [Martin et al., 2003], thus providing information on the 'smearing' effect due to transport (significant in fall and winter with longer NOx lifetime) in the latter approach. III. Methodology

### **SCIAMACHY NO<sub>2</sub> Retrieval**

>On board ESA/ENVISAT satellite launched in March 2002. -Sun-synchronous orbit, crossing equator 10:00 AM local time in the descending node. -Nadir view with resolution of 30 km x 60 km (6days for global coverage).

(1) Obtaining total slant columns of  $NO_2$  by direct fitting backscattered radiance spectra (429 – 452nm) with reference spectra [*Chance*, 1998; *Martin et al.*, 2002].

(2) Removing stratospheric columns by subtracting central Pacific NO2 columns from total columns [*Martin et al.*, 2002]  $\rightarrow$  Tropospheric NO<sub>2</sub> slant columns.

(3) Calculating air mass factor (AMF) to convert tropopospheric NO<sub>2</sub> slant columns into vertical tropospheric NO2 columns [*Palmer et al.,* 2001; *Spurr et al.,* 2001, 2002]. \* Integral of the relative vertical distrution of NO<sub>2</sub> (shape factor)  $\leftarrow$  GEOS-Chem data. \* Shape factor is weighted by local sensitivity of backscattered radiance to  $NO_2$ (scattering weights). ← Linearized Discrete Ordinate Radiative Transfer (LIDORT) model.

(4) Cloud correction with Fast Retrieval Scheme for Cloud Observables (FRESCO)

# Adjoint Inversion of Global NO<sub>x</sub> emissions with SCIAMACHY NO<sub>2</sub>

Changsub Shim<sup>1</sup>, Qinbin Li<sup>1</sup>, Daven Henze<sup>2</sup>, Randall V. Martin<sup>3</sup>, Aaron Daankellar<sup>3</sup>, Monika Kopacz<sup>4</sup>, and Kevin Bowman<sup>1</sup> <sup>1</sup>Jet Propulsion Laboratory, California Institute of Technology, Pasadena, CA <sup>2</sup> NASA Goddard Institute of Space Sciences, Columbia University, NYC, NY <sup>3</sup> Department of Physics and Atmospheric Science, Dalhousie University, Halifax, Canada <sup>4</sup> Division of Engineering and Applied

Sciences, Harvard University, Cambridge, MA. Email: cshim@jpl.nasa.gov



<sup>0.18 0.54 0.89 1.25 1.61 1.96 2.32 2.68 3.04 3.39 3.93 5.00</sup> 

 $\succ$ Cost function reduction: ~50% after 6 iterations. Some improvements over the overestimated regions (Blue) Little improvements over the underestimated regions (Red)

$$\mathbf{K} = \frac{\partial \mathbf{y}_{\mathbf{n}}}{\partial \mathbf{x}} = \frac{\partial \mathbf{y}_{\mathbf{n}}}{\partial \mathbf{y}_{\mathbf{n-1}}} \frac{\partial \mathbf{y}_{\mathbf{n-1}}}{\partial \mathbf{y}_{\mathbf{n-2}}} \dots \frac{\partial \mathbf{y}_{\mathbf{n}}}{\partial \mathbf{y}_{\mathbf{0}}} \frac{\partial \mathbf{y}_{\mathbf{0}}}{\partial \mathbf{x}}$$

$\mathbf{K} = \left(\frac{\overline{\partial \mathbf{y}_{n-1}}}{\overline{\partial \mathbf{y}_{n-2}}} \cdots \overline{\partial \mathbf{y}_{0}} \frac{\overline{\partial \mathbf{x}}}{\overline{\partial \mathbf{x}}}\right) = \left(\frac{\overline{\partial \mathbf{x}}}{\overline{\partial \mathbf{x}}}\right) \left(\frac{\overline{\partial \mathbf{y}_{0}}}{\overline{\partial \mathbf{y}_{0}}}\right) \cdots \left(\frac{\overline{\partial \mathbf{y}_{n-2}}}{\overline{\partial \mathbf{y}_{n-2}}}\right) \left(\frac{\overline{\partial \mathbf{y}_{n-1}}}{\overline{\partial \mathbf{y}_{n-1}}}\right)$	$\mathbf{K}^{T} = $	$\left(\frac{\partial y_n}{\partial y_{n-1}}\frac{\partial y_{n-1}}{\partial y_{n-2}}\right)$	$\left(\frac{\partial \mathbf{y}_1}{\partial \mathbf{y}_0} \frac{\partial \mathbf{y}_0}{\partial \mathbf{x}}\right)^T$	$= \left(\frac{\partial \mathbf{y}_0}{\partial \mathbf{x}}\right)^T \left(\frac{\partial \mathbf{y}_1}{\partial \mathbf{y}_0}\right)^T$	$\dots \left(\frac{\partial \mathbf{y}_{\mathbf{n}-1}}{\partial \mathbf{y}_{\mathbf{n}-2}}\right)^T \left(\frac{\partial \mathbf{y}_{\mathbf{n}}}{\partial \mathbf{y}_{\mathbf{n}-1}}\right)$	
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SCIAMACHY – a posteriori CTM



-4.00-2.00-0.90-0.70-0.50-0.30-0.10 0.20 0.60 1.00 3.00 5.00 10<sup>15</sup> molecules/cm<sup>2</sup>



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	, e. e. p.:.e.,		
/yr	%		
24)	31		
·9)	18.7		
· <b>6</b> )	31.5		
7)	12.5		
	~5		
	<1		
9 1	<1		
	<1		
3)	100		
nd Davidson et al. (1991			